

Measuring and modelling broadband albedo on Lewis Glacier, Mt Kenya



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Introduction

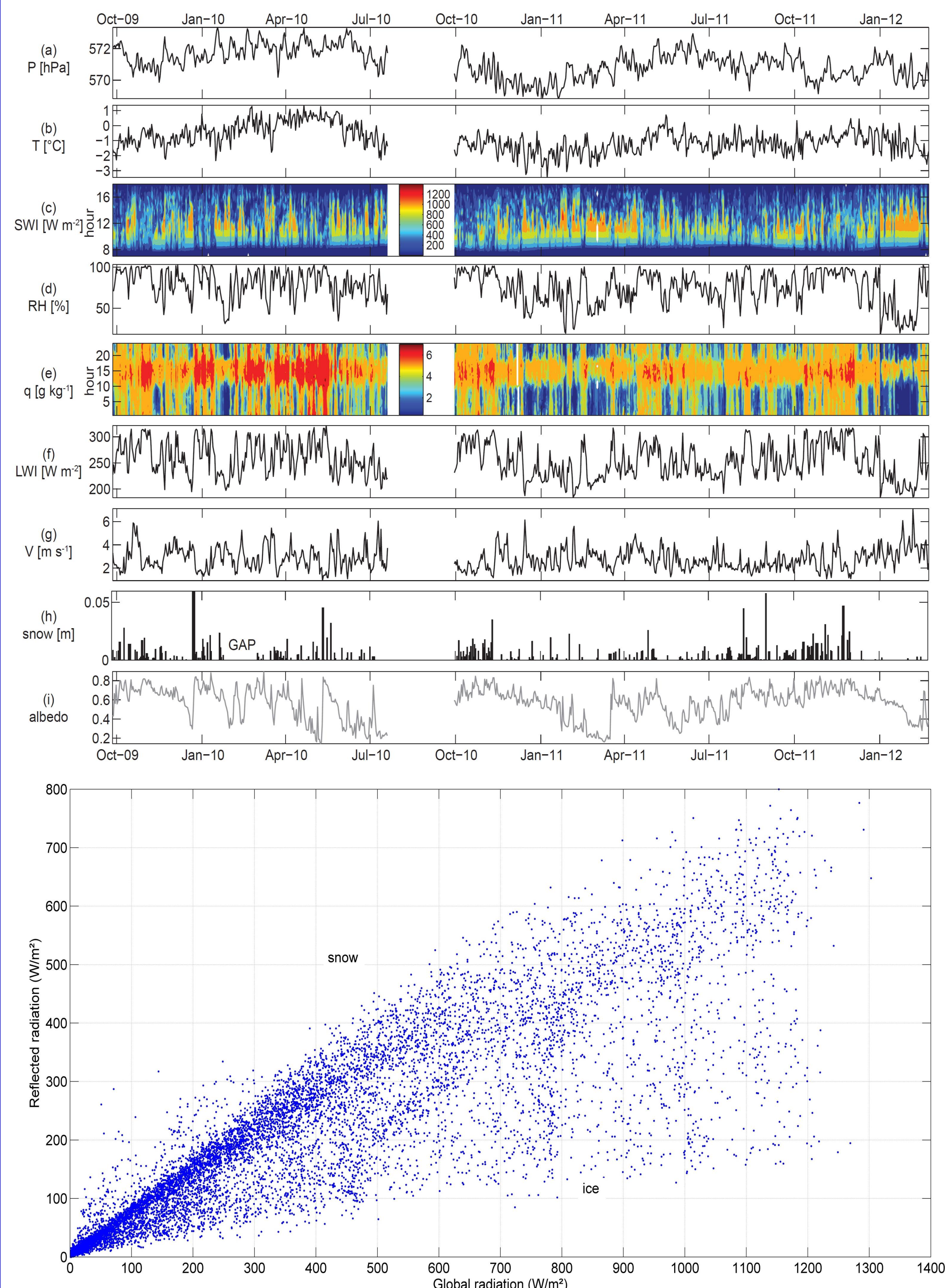
Footprint of the tropical (0.15° S) and high altitude (4830 m) climate

- Hygric seasons and high diurnal climate variability
- Snow fall events not confined to seasons and occurring almost daily
- High diurnal variability in energy balance: fresh snow in the afternoon, fast decay of snow albedo in the morning, melt around noon
- High sublimation rates
- Fresh snow events frequently at the limit of the measurement accuracy (1 cm)

Challenge: Calibrate albedo parameterization for distributed mass and energy balance modelling!

Measurements

In-situ observations of different climate parameters (Nicholson et al. 2013) and incoming versus reflected short-wave radiation (scatter plot) performed on Lewis Glacier between 01 Oct 2009 and 22 Feb 2012. Radiative fluxes were recorded by a Kipp&Zonen CNR4 net radiometer above the 15° inclined glacier surface.



Modelling

The parameterization of albedo employs an approach from Oerlemans and Knap (1998), where the albedo (α) is a function of the snowfall frequency, i.e. time since the last snowfall ($s-i$), and snow depth (d). The effects of snow ageing (t^*) and snow depth on the albedo are given by e-folding constants to compute albedo values between constant fresh snow and firn albedos and variable ice albedo, which is a linear function of the dew point temperature (DPT) to account for variations of the ice albedo due to sublimation (Mölg et al. 2009).

$$\alpha_{snow} = \alpha_{firn} + (\alpha_{frsnow} - \alpha_{firn}) \exp\left(\frac{s-i}{t^*}\right) \quad (1)$$

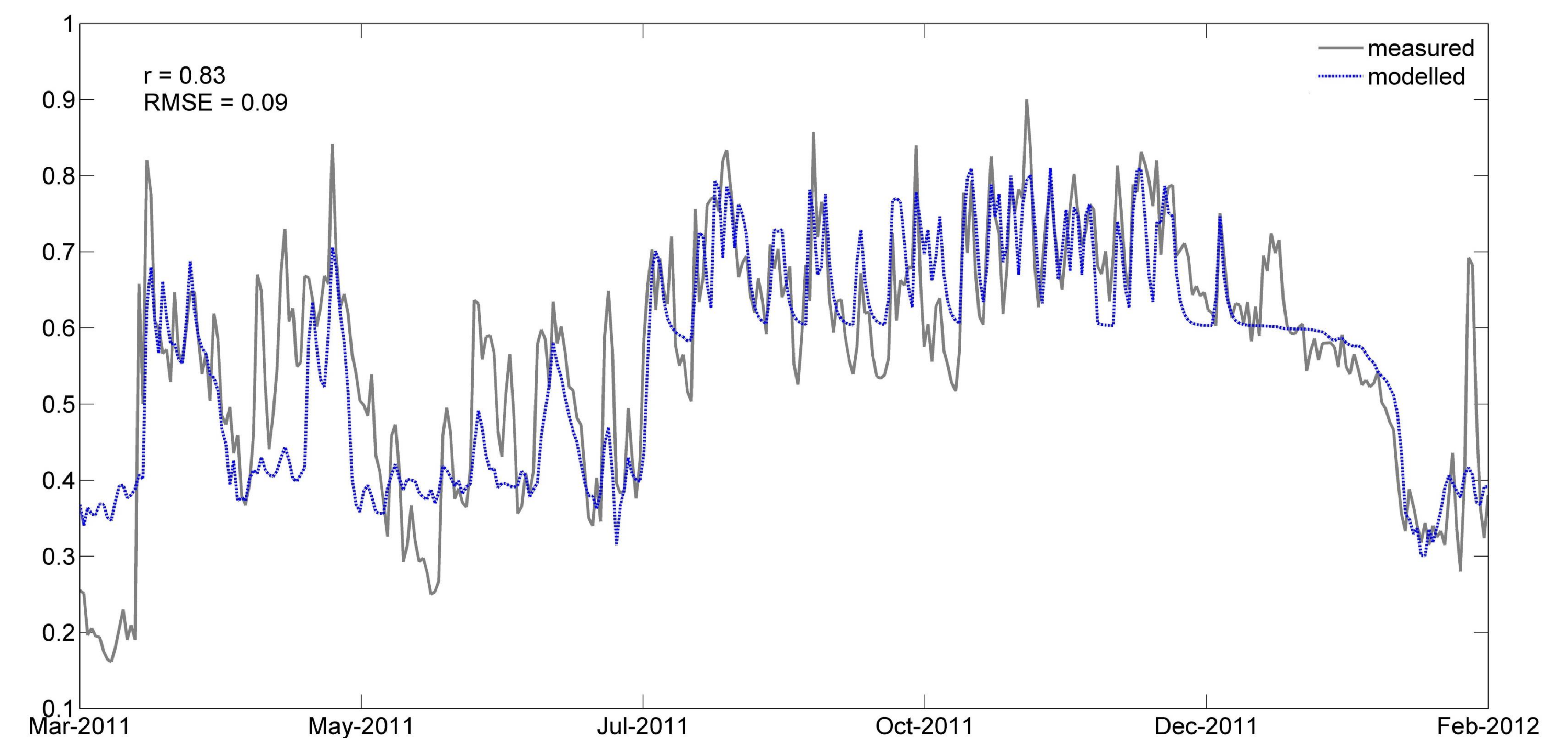
$$\alpha = \alpha_{snow} + \alpha_{ice} - \alpha_{snow} \exp\left(\frac{-d}{d^*}\right) \quad (2)$$

$$\alpha_{ice} = c_1 DPT + c_2 \quad (3)$$

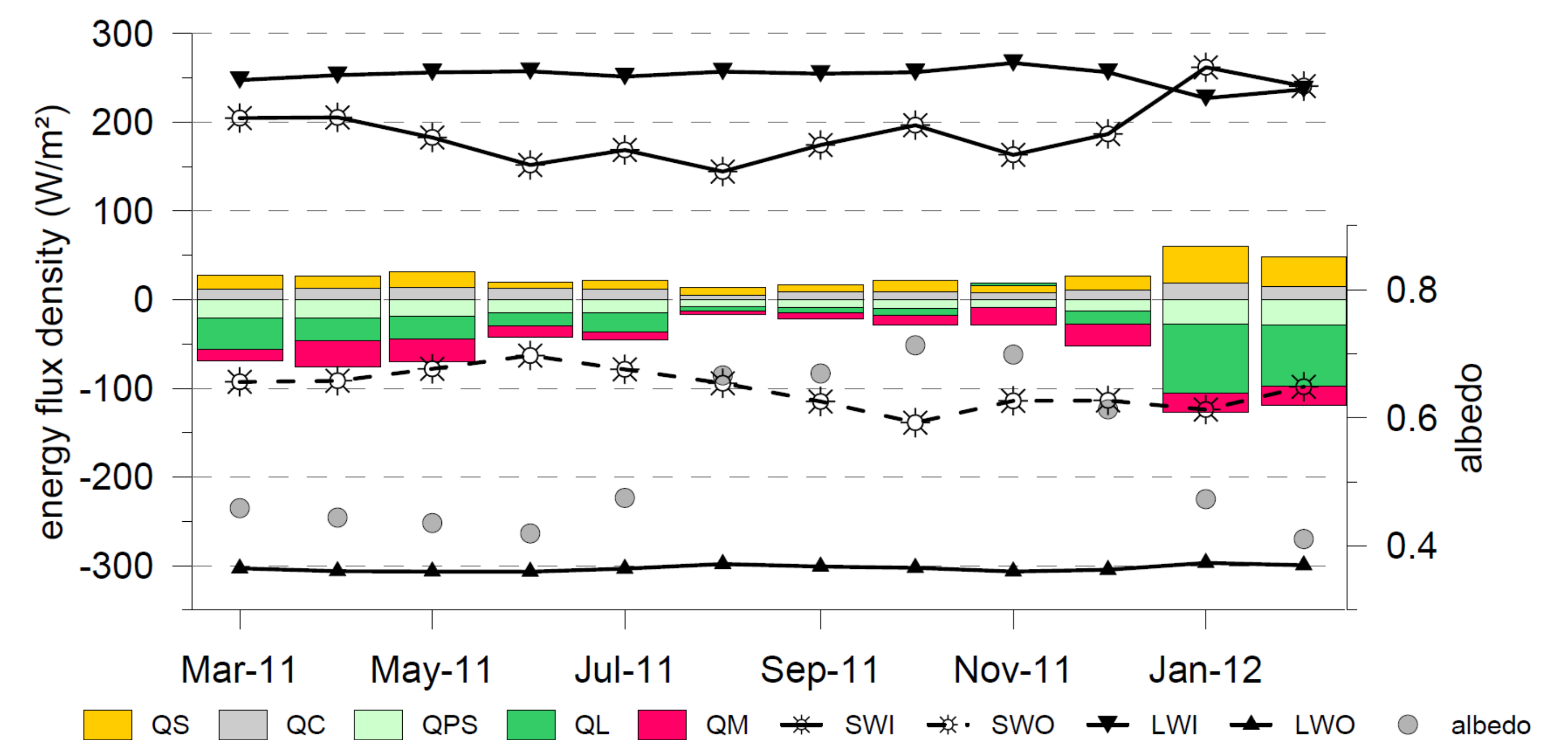
Results

Optimal values were found by a Monte Carlo simulation, including the total mass balance model parameter setup, using a subsequent cross validation confirming the best parameter setup (Prinz et al. 2016).

$$\begin{aligned} \alpha_{firn} &= 0.60, & \alpha_{frsnow} &= 0.81 \\ t^* &= 1.3 \text{ (days)}, & d^* &= 4.1 \text{ (cm)} \\ c_1 &= 0.0056 \text{ (}^\circ\text{C}^{-1}\text{)}, & c_2 &= 0.4179 \end{aligned}$$



Measured and modelled mean daily albedo using equations and values given above.



Resulting mean monthly energy flux densities of the final, distributed model setup for the mass balance year 2011/12.

References:

- Mölg, T., Cullen, N.J., Hardy, D.R., Kaser, G. and Klok, L. (2008): Mass balance of a slope glacier on Kilimanjaro and its sensitivity to climate. *Int. J. Climatol.*, 28, 881-892, doi: 10.1002/joc.1589.
- Nicholson, L.I., Prinz, R., Mölg, T., Kaser, G., (2013): Micrometeorological conditions and surface mass and energy fluxes on Lewis Glacier, Mt Kenya, in relation to other tropical glaciers. *The Cryosphere* 7, 1205–1225. doi:10.5194/tc-7-1205-2013
- Oerlemans, J., and W. H. Knap (1998): A 1 year record of global radiation and albedo in the ablation zone of Morteratschgletscher, Switzerland, *Journal of Glaciology*, 44(147), 231–238.
- Prinz, R., Nicholson, L., Mölg, T., Gurgiser, W. and Kaser, G. (2016): Climatic controls and climate proxy potential of Lewis Glacier, Mt Kenya, *The Cryosphere* 10, 133-148, doi:10.5194/tc-10-133-2016.